

# Understanding ENSO dynamics through the exploration of past climates

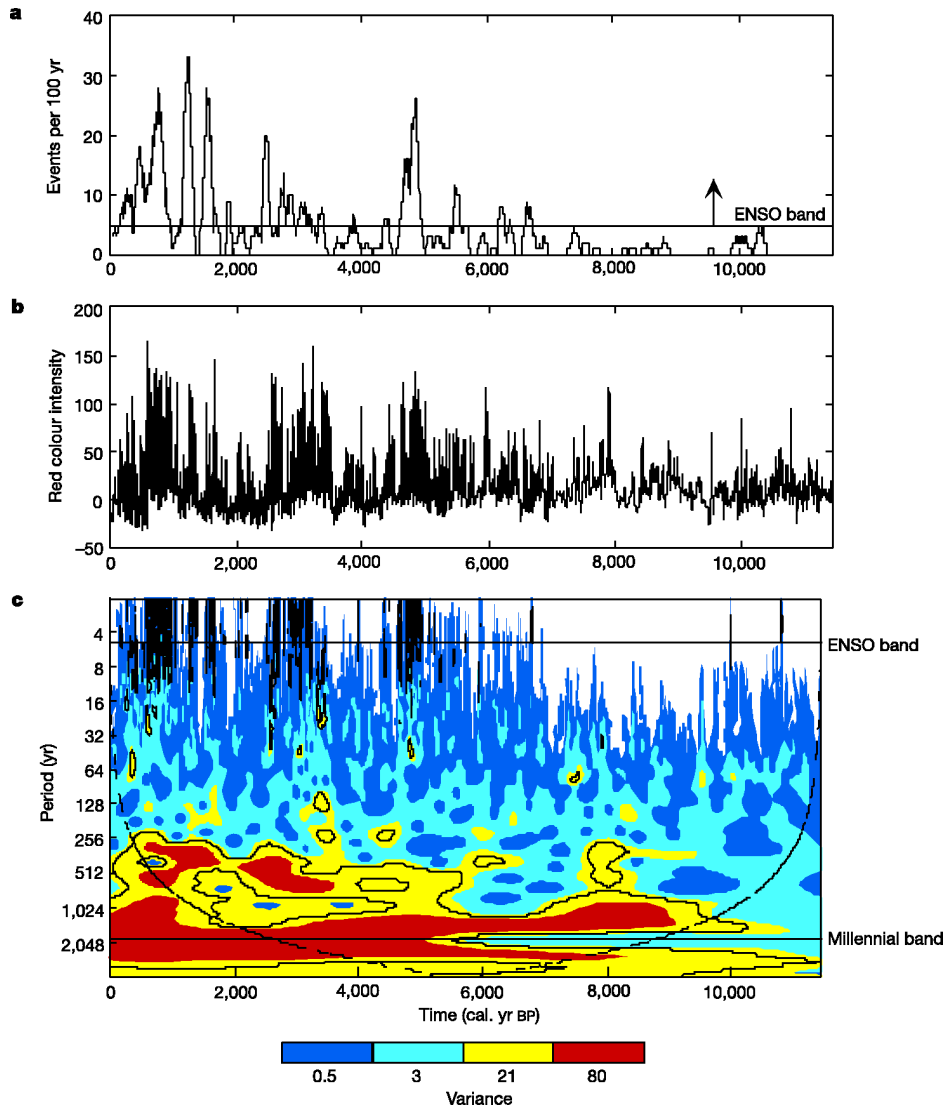
Steven J. Phipps<sup>1</sup>

Jaclyn N. Brown<sup>2</sup>

<sup>1</sup>Climate Change Research Centre, UNSW, Sydney, Australia

<sup>2</sup>Centre for Australian Weather and Climate Research, Hobart, Australia

# El Niño has changed...

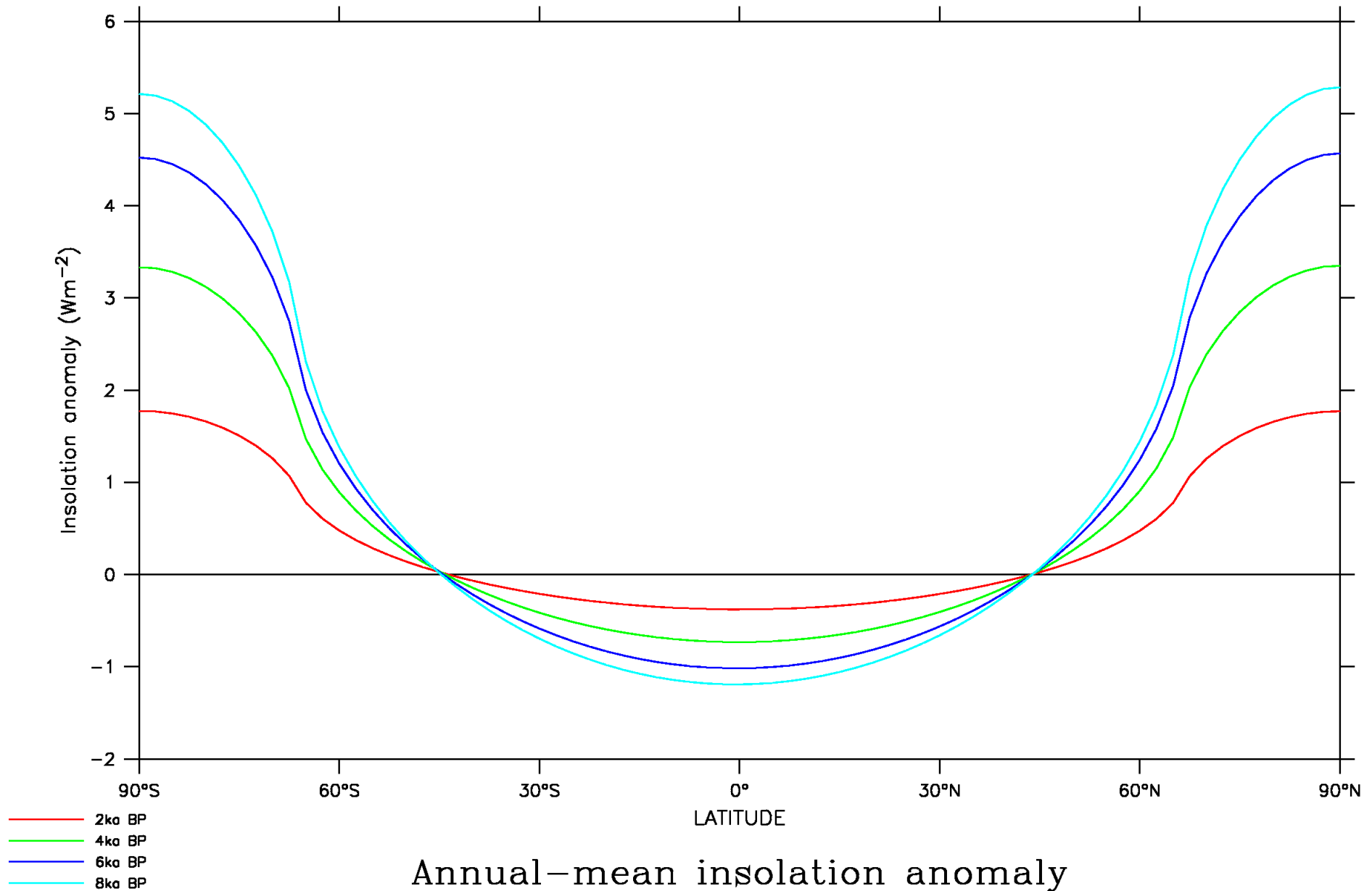


Moy et al. (2002), *Nature*

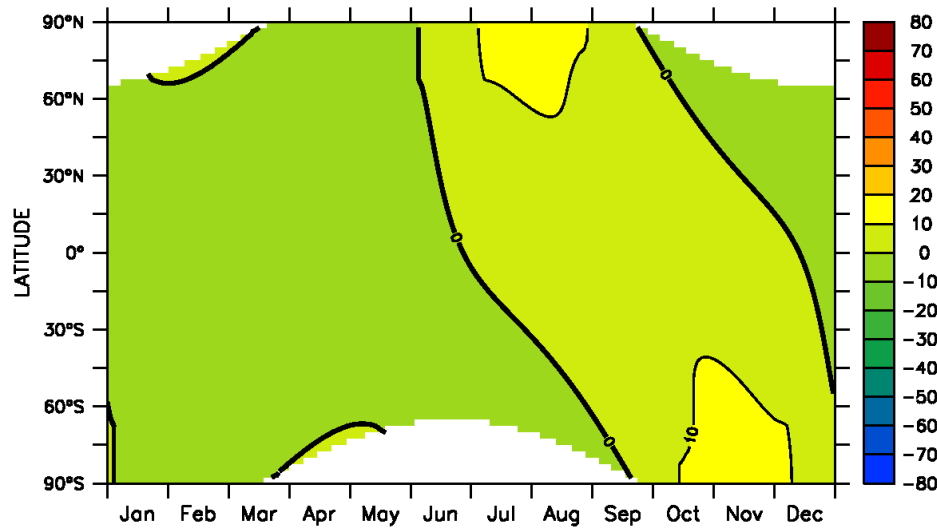
- “Modern” El Niño began 7-5 ka BP, with only weak decadal-scale events beforehand
- El Niño was 15-60% weaker at 6 ka BP than at present
- Gradual strengthening of El Niño thereafter
- Evidence of a peak in strength at 2-1 ka, possibly earlier in the western Pacific than in the east



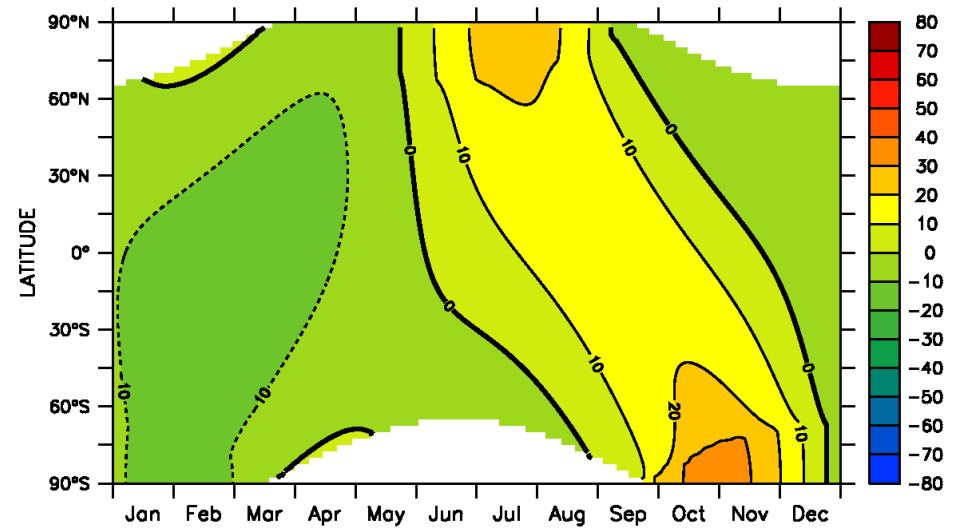
# The changes in annual-mean insolation are small...



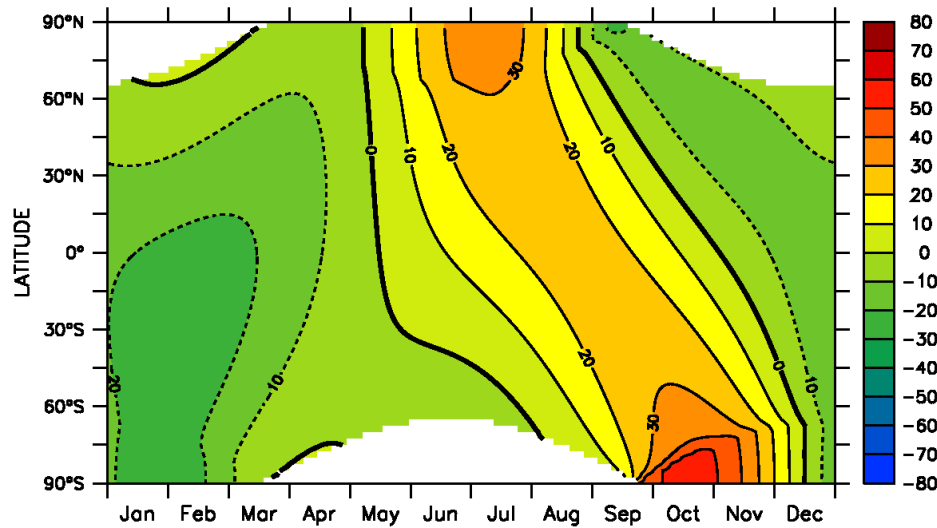
... but the seasonal changes are large



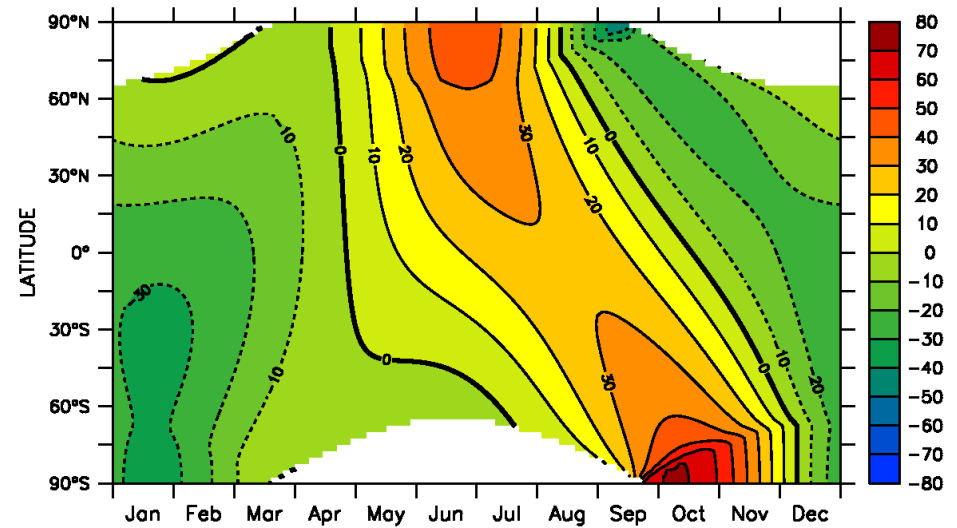
Insolation anomaly ( $\text{Wm}^{-2}$ ): 2ka BP



Insolation anomaly ( $\text{Wm}^{-2}$ ): 4ka BP



Insolation anomaly ( $\text{Wm}^{-2}$ ): 6ka BP



Insolation anomaly ( $\text{Wm}^{-2}$ ): 8ka BP

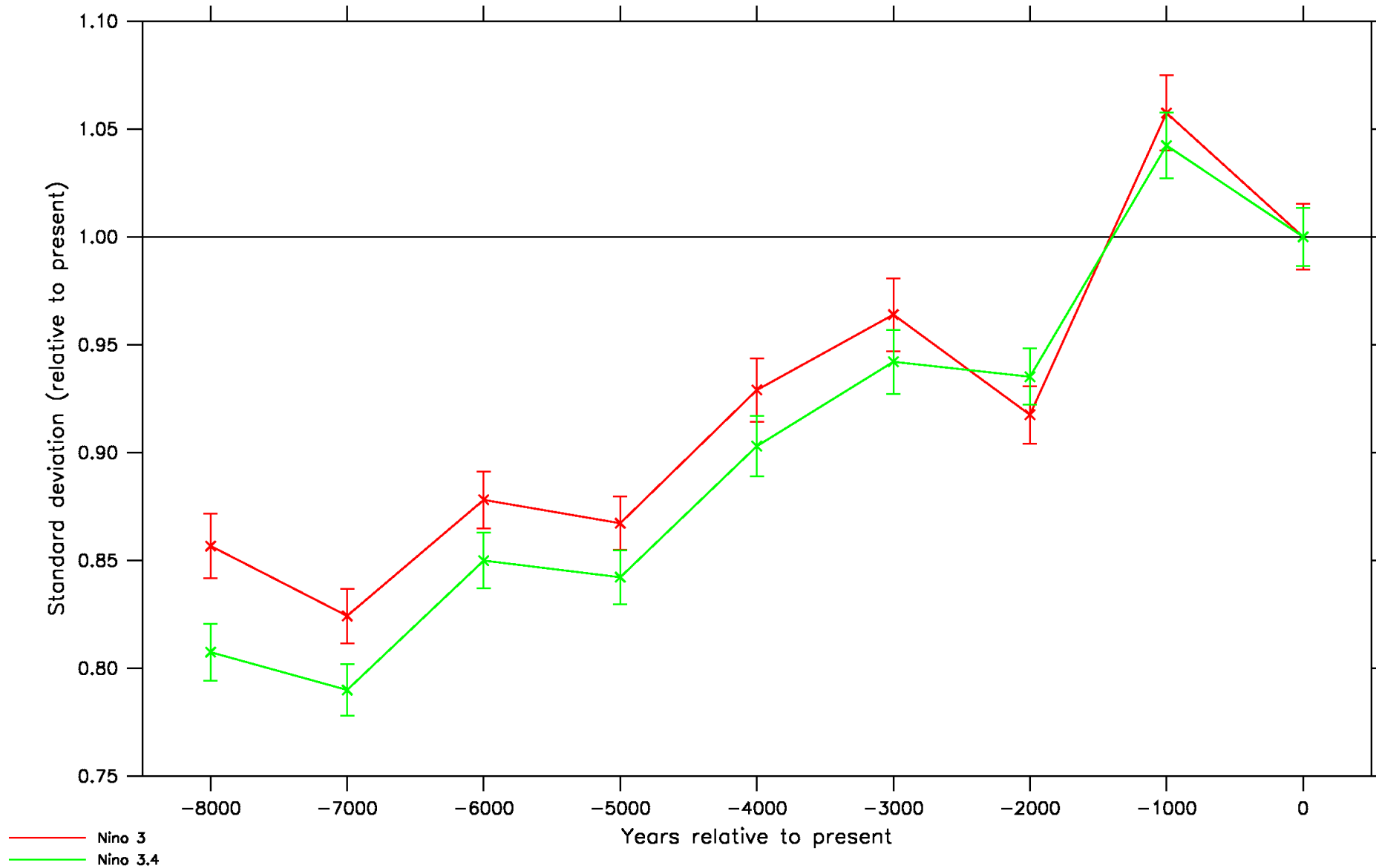
# Current understanding

- Previous modelling work has shown that orbitally-driven changes in insolation can alter ENSO behaviour
- Broadly consistent mechanism found to explain weaker mid-Holocene ENSO:
  - Insolation changes result in enhanced seasonal cycle in NH
  - Intensification of summer monsoon system
  - Enhanced Walker circulation
  - Stronger easterly trade winds in central and western Pacific
  - Steeper thermocline/increased upwelling in central and eastern Pacific
  - Suppresses development of El Niño events
- However, this proposed mechanism is qualitative in nature and has yet to be rigorously tested

# Simulations of the late Holocene climate

- CSIRO Mk3L climate system model v1.1:
  - Atmosphere: R21 ( $5.6^\circ \times 3.2^\circ$ ), 18 vertical levels
  - Ocean:  $2.8^\circ \times 1.6^\circ$ , 21 vertical levels
  - Sea ice: Dynamic-thermodynamic
  - Land surface: Static vegetation
  - Flux adjustments applied
- Snapshot simulations for 8, 7, 6, 5, 4, 3, 2, 1 and 0 ka BP:
  - Only the Earth's orbital parameters are varied
  - Atmospheric CO<sub>2</sub> concentration = 280ppm
  - Solar constant =  $1365 \text{ Wm}^{-2}$
  - Integrated for 1000 years

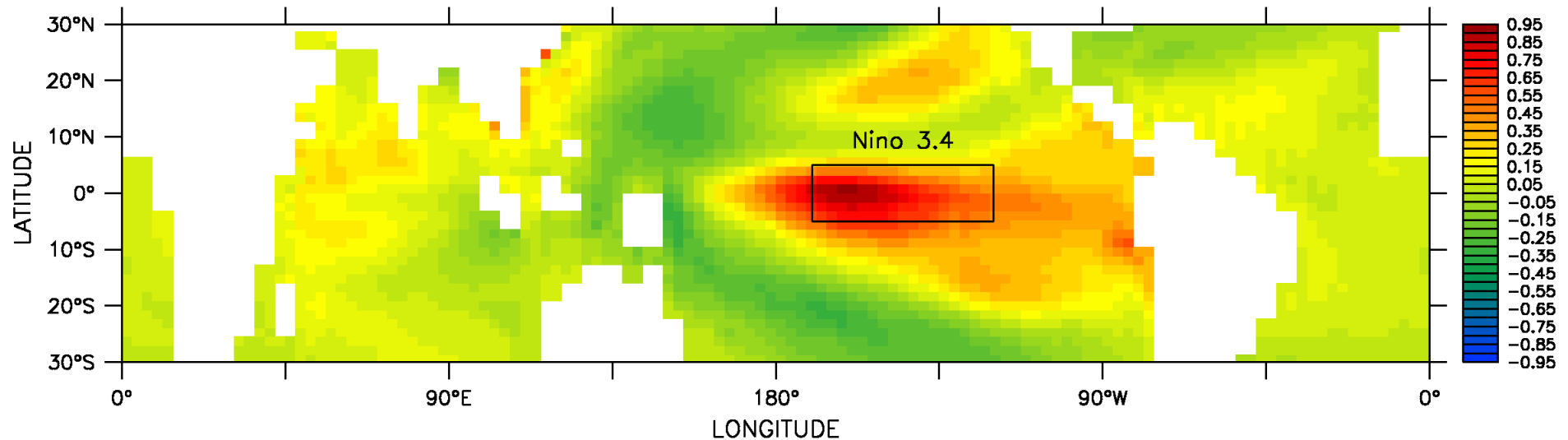
# Simulated changes in ENSO variability



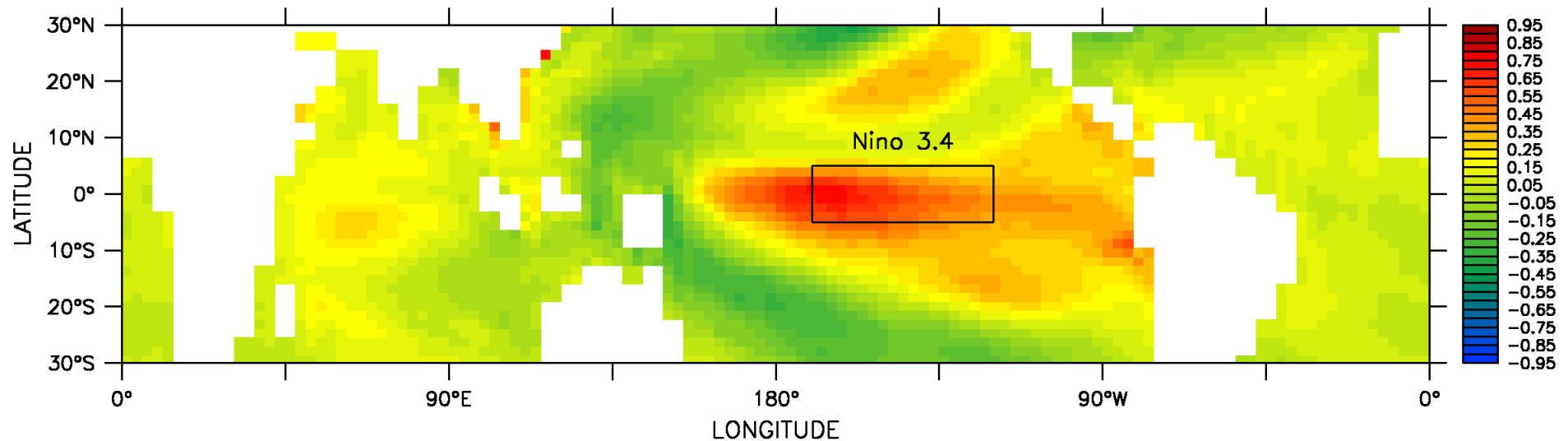
Standard deviation of Nino SST anomaly



# ENSO strengthens and shifts eastwards

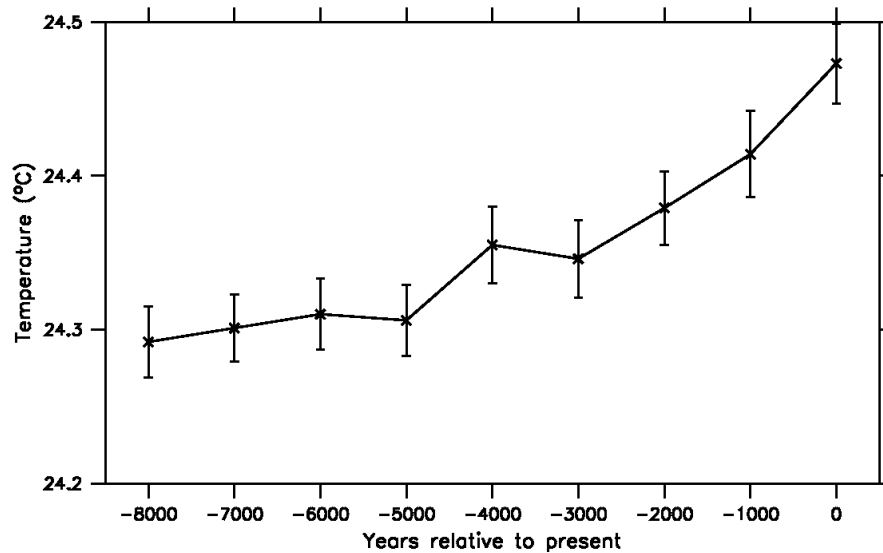


EOF1 of monthly SST anomalies (°C): 0ka BP

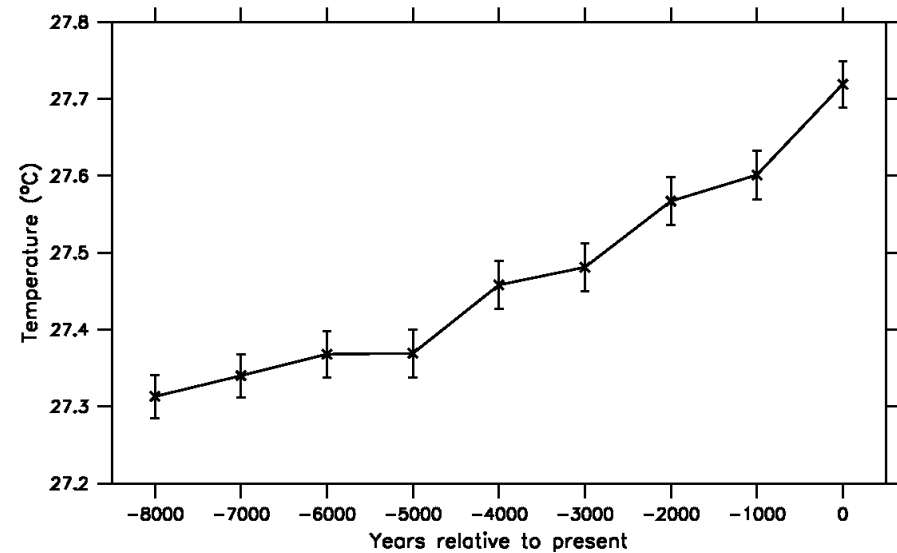


EOF1 of monthly SST anomalies (°C): 8ka BP

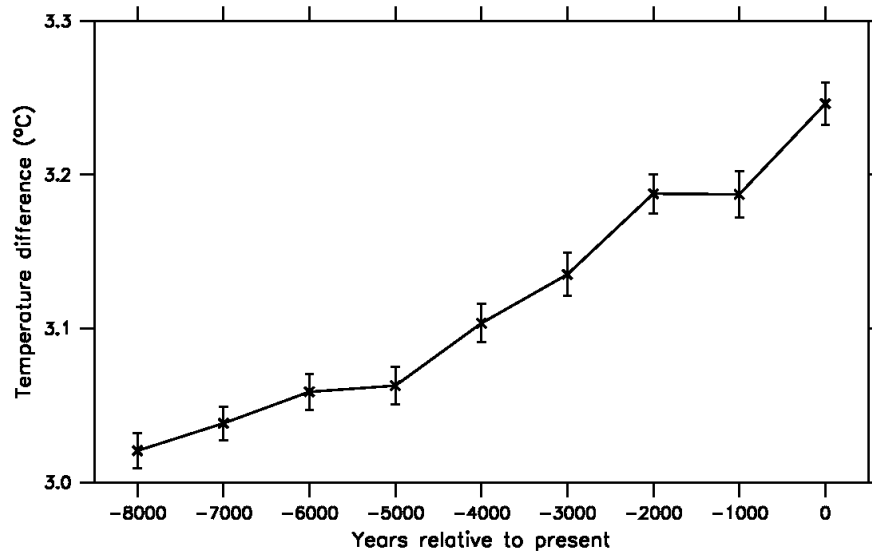
# The changes in the mean state are small...



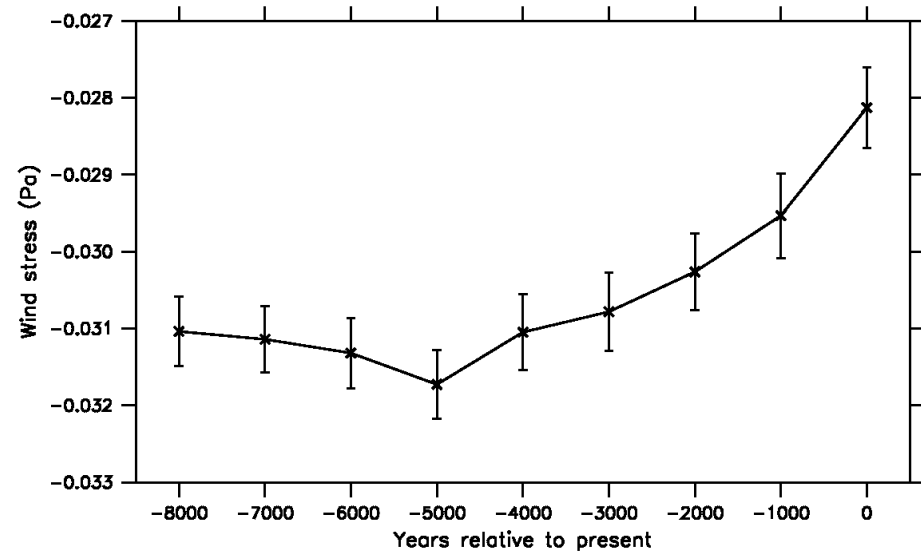
(a) Annual-mean Nino 3 SST



(b) Annual-mean Nino 4 SST

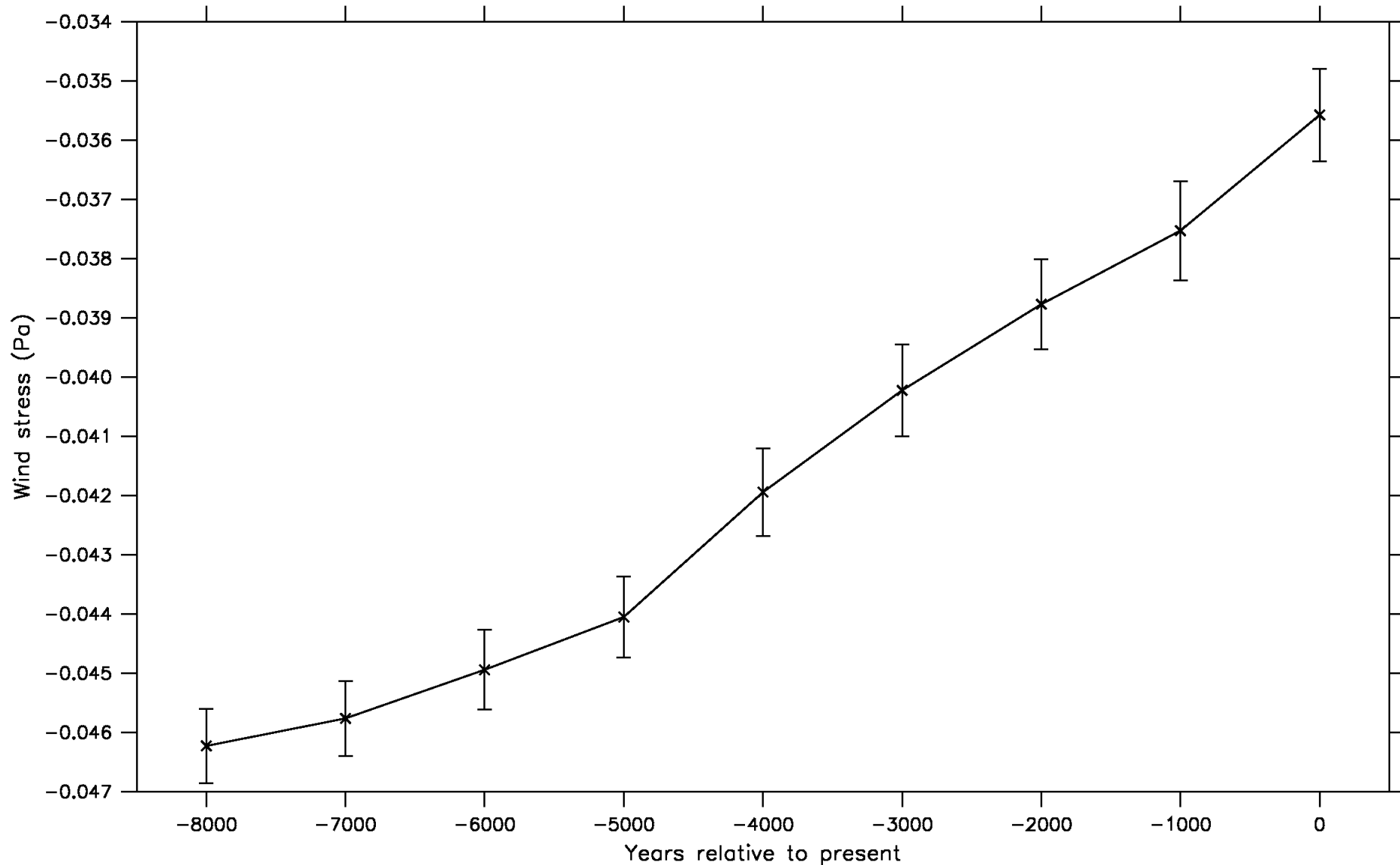


(c) Nino 4 SST minus Nino 3 SST



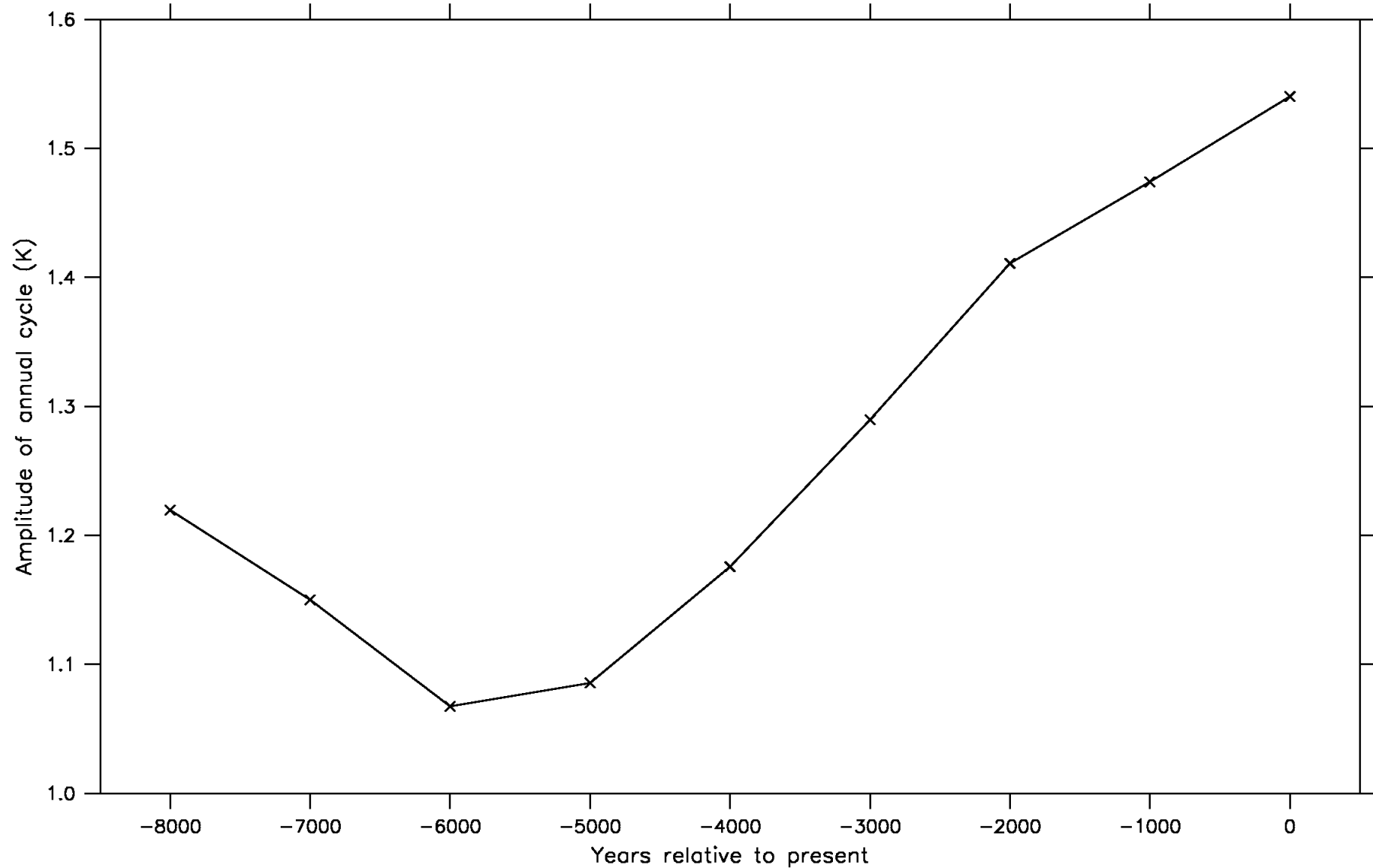
(d) Annual-mean Nino 4 zonal wind stress

... but the seasonal changes are larger



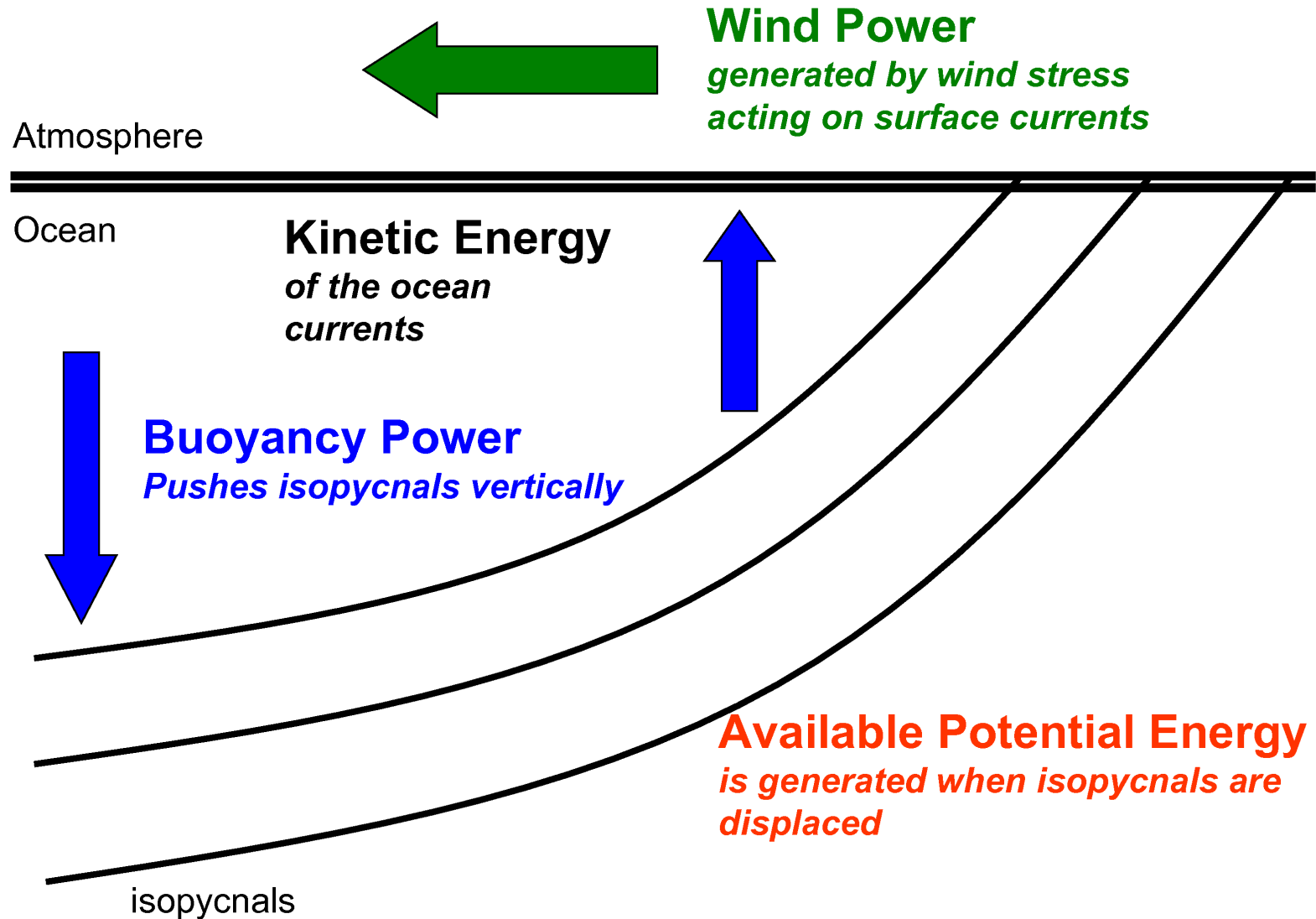
JASO zonal wind stress in Niño 4 region

# No apparent correlation with the annual cycle

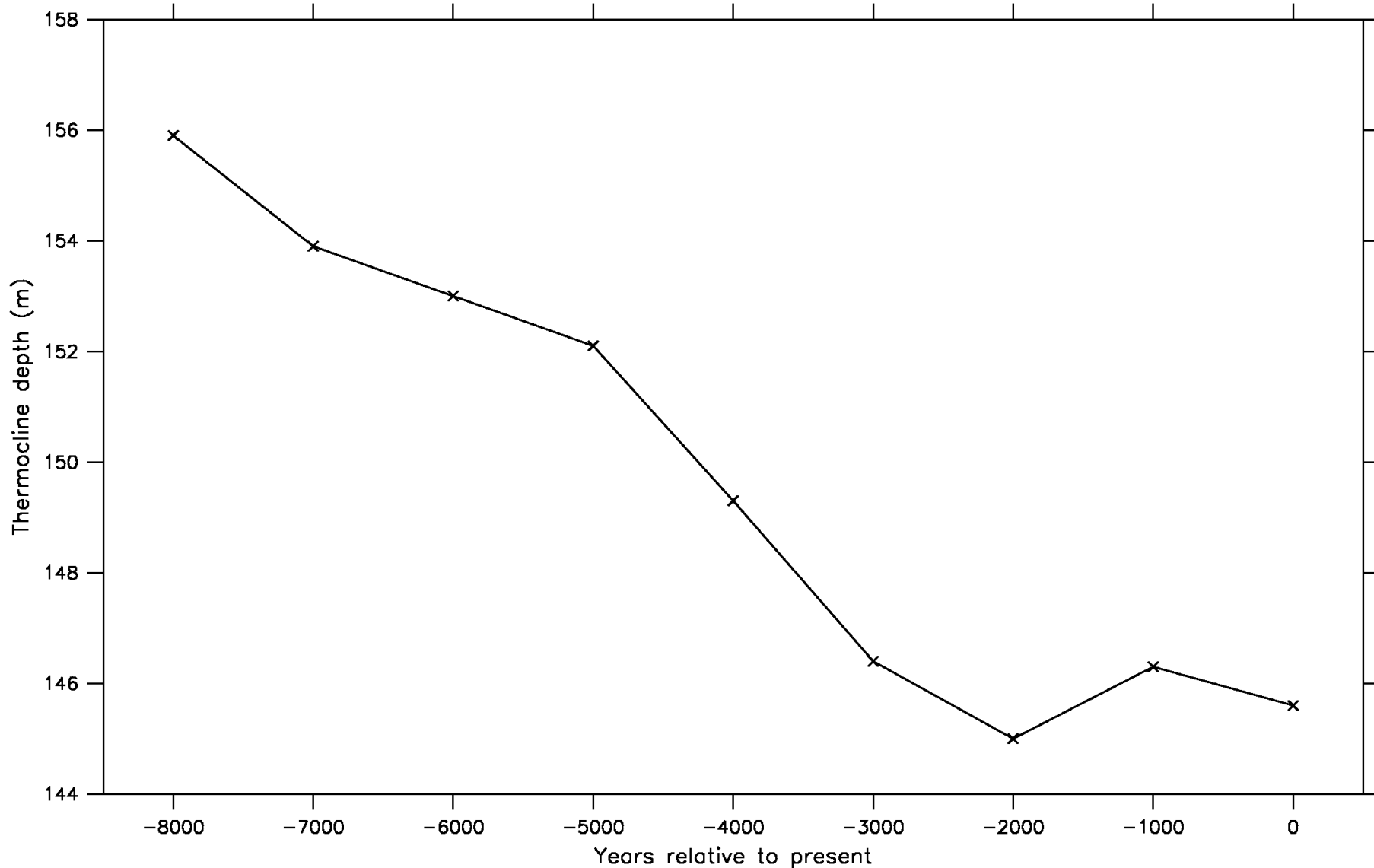


Amplitude of annual cycle in Nino 3 SST

# Ocean energetics



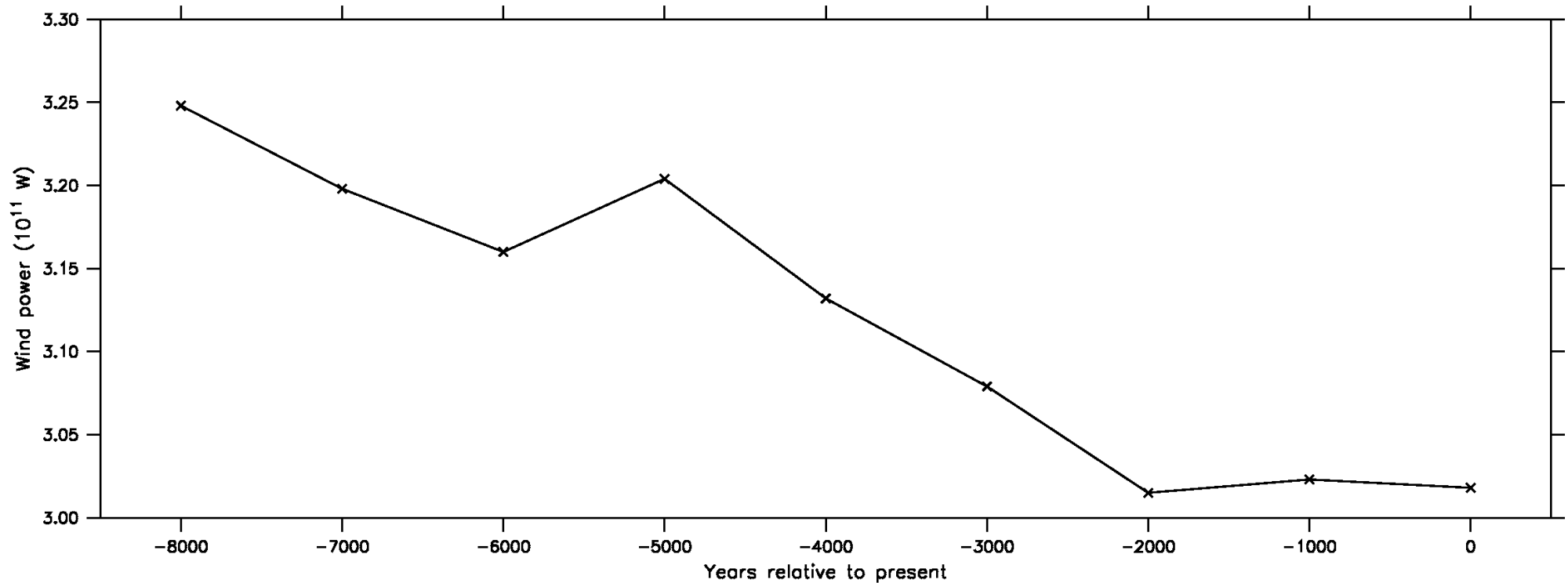
# The thermocline becomes shallower



Mean thermocline depth in equatorial Pacific

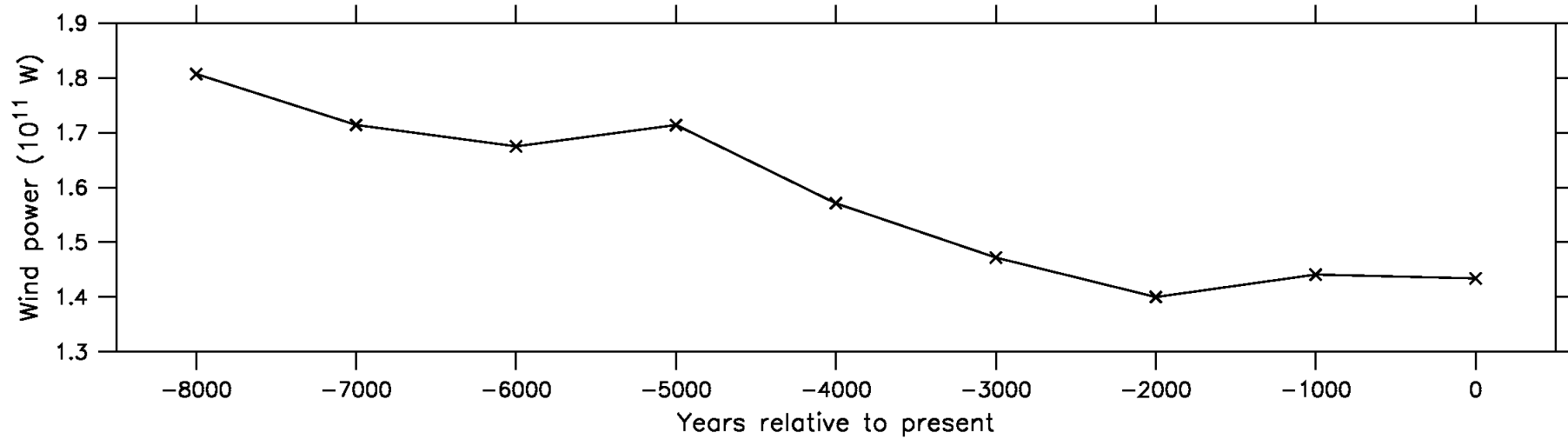
# The changes in annual-mean wind power are small...

$$W = \iint_{z=0} \underline{u} \cdot \underline{\tau} \, dx dy$$

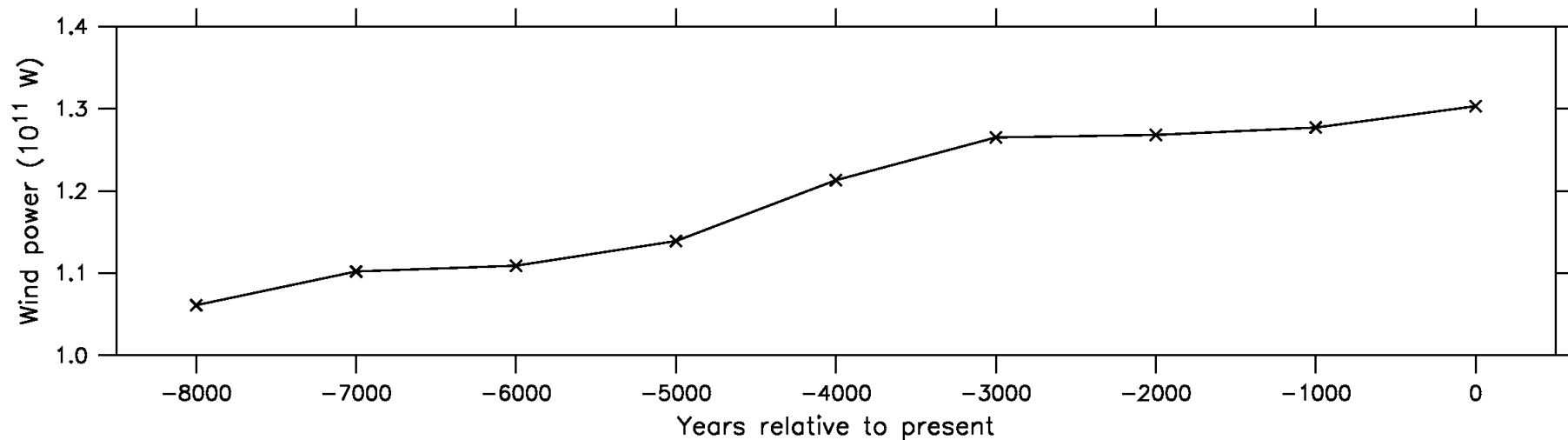


Annual-mean wind power

... but the seasonal changes are larger



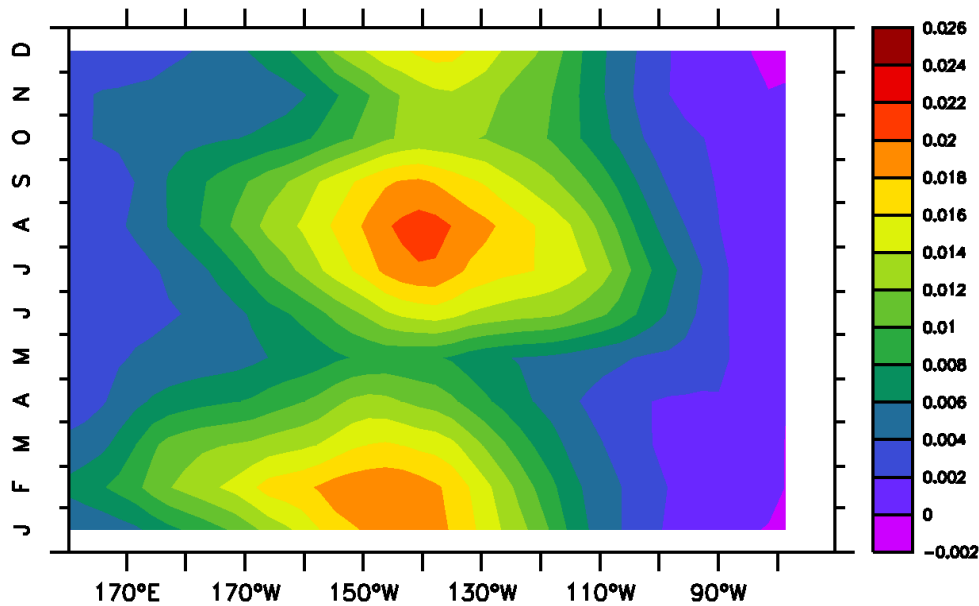
June–July–August–September wind power



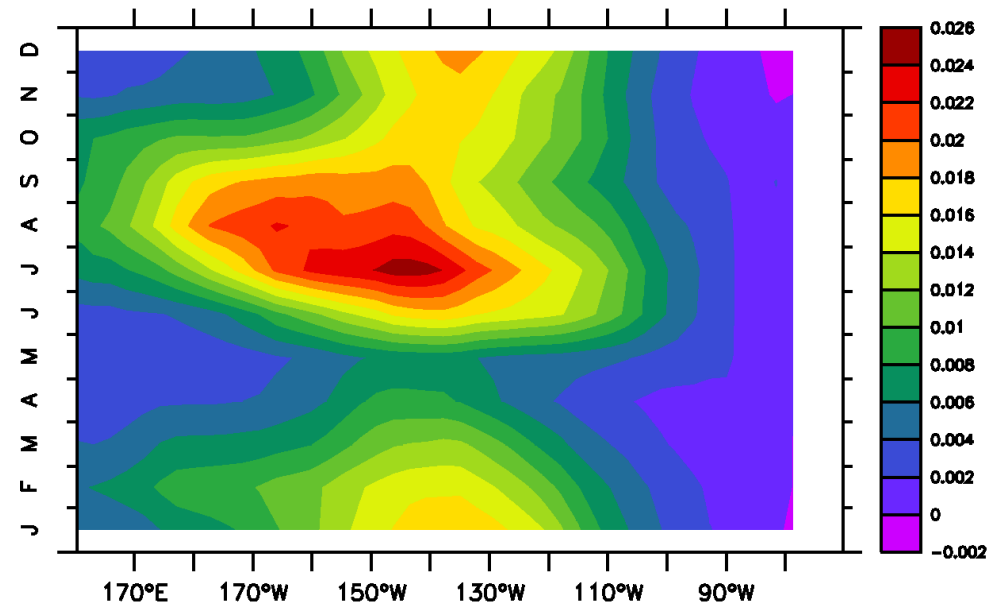
January–February–March–April wind power



# Annual cycle in wind power on the equator



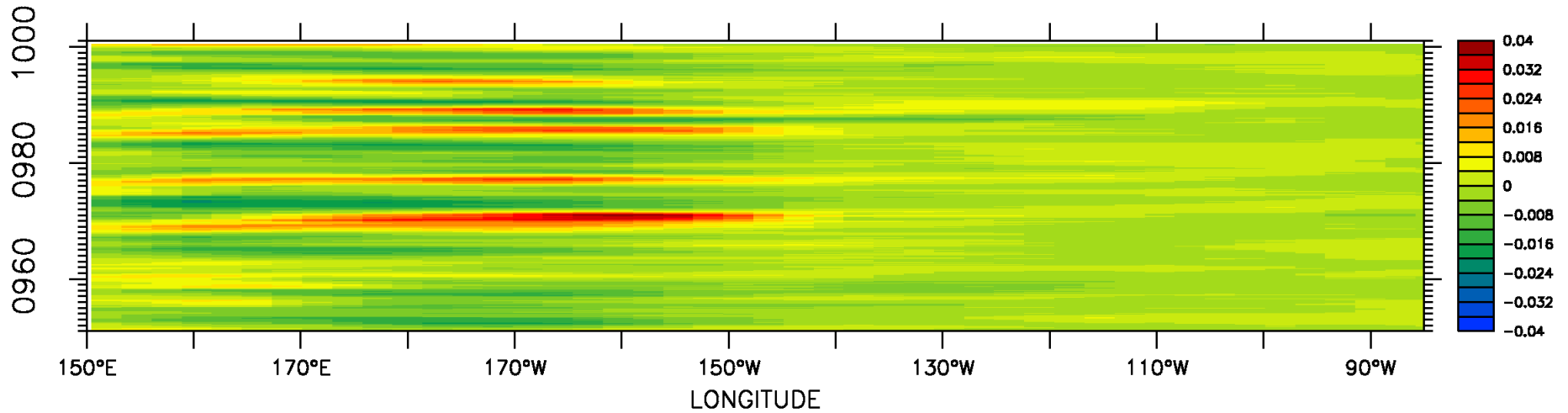
0 ka



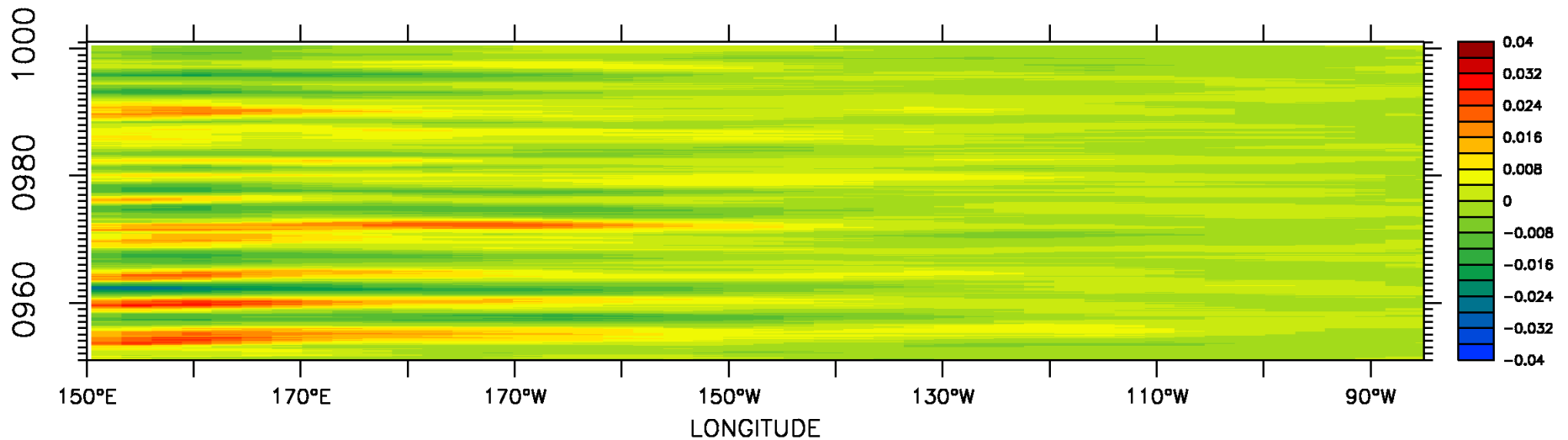
8 ka

Wind power ( $\text{Wm}^{-2}$ )

# Propagation of Westerly Wind Bursts



Zonal wind stress anomaly at equator (Pa): 0ka BP



Zonal wind stress anomaly at equator (Pa): 8ka BP

# Conclusions

- The study of palaeo-ENSO allows us to explore the links between ENSO and the global climate system.
- By forcing a model with orbitally-driven insolation changes only, we are able to broadly reproduce the changes in ENSO behaviour over the Holocene.
- Physical links between ENSO, the Walker Circulation and the Asian monsoon appear to explain the upward trend in variability.
- However, it does not explain the peak at 1 ka. Other mechanisms therefore appear to be at work.
- A full understanding of the processes that drive changes in ENSO variability may be within grasp. However, this will require an approach that integrates the theory, data and modelling communities.