



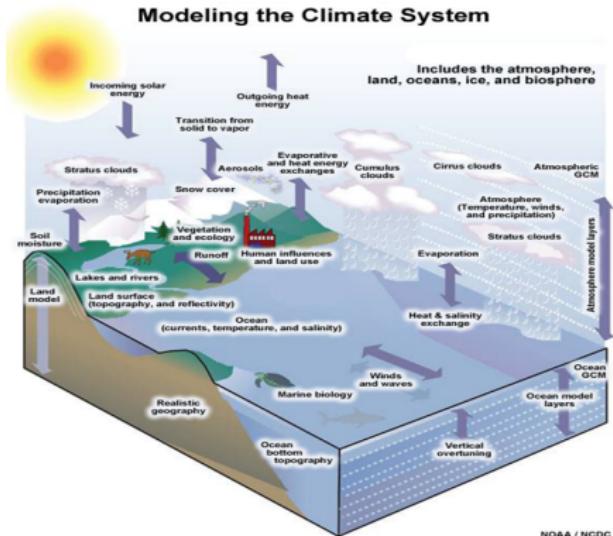
Sensitivity of the Antarctic ice sheet to marine climate variability and change

Steven J. Phipps
Institute for Marine and Antarctic Studies
University of Tasmania

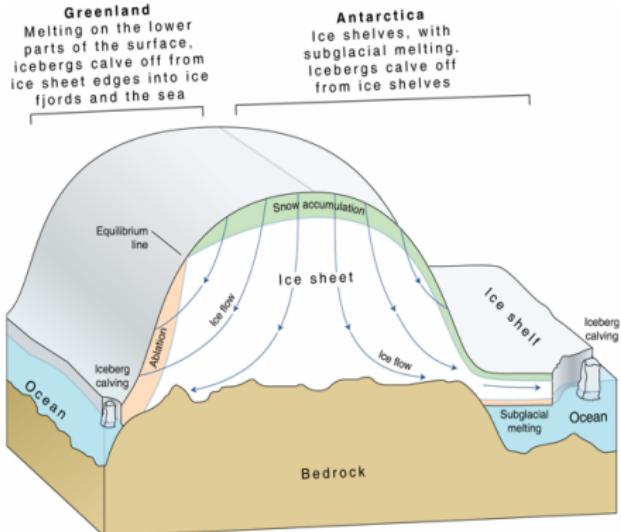
Elizabeth and Frederick White Conference
5–7 July 2017

Climate and ice sheet modelling

Climate model



Ice sheet model



Images courtesy of NOAA/NCDC and NSIDC

Differing degrees of complexity

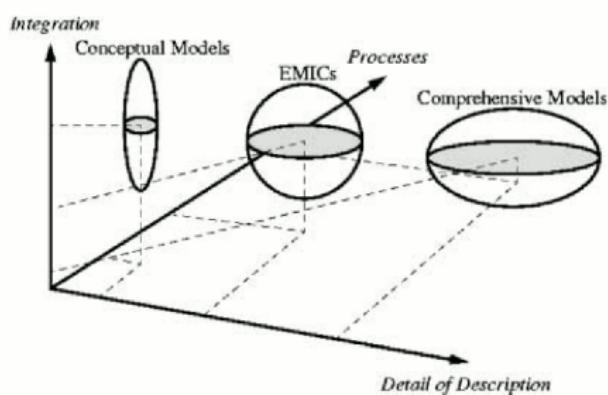


Fig. 1. Pictorial definition of EMICs. Adapted from Claussen (2000)

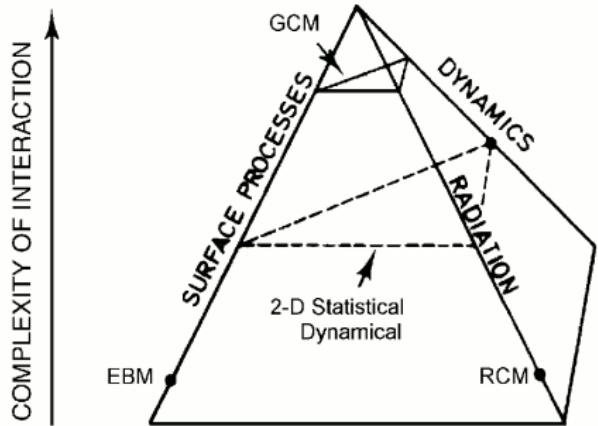
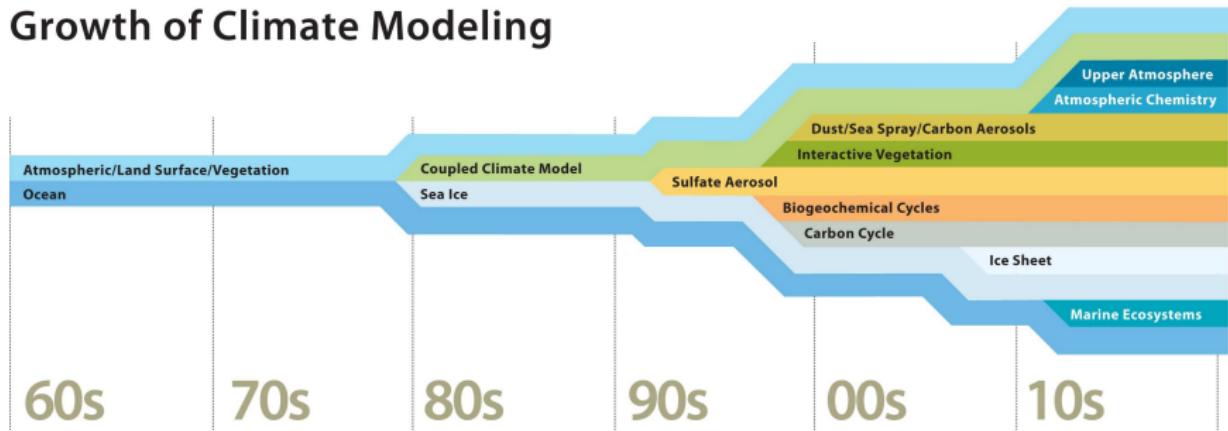


Fig. 2. The climate modeling pyramid. Adapted from Henderson-Sellers and McGuffie (1987)

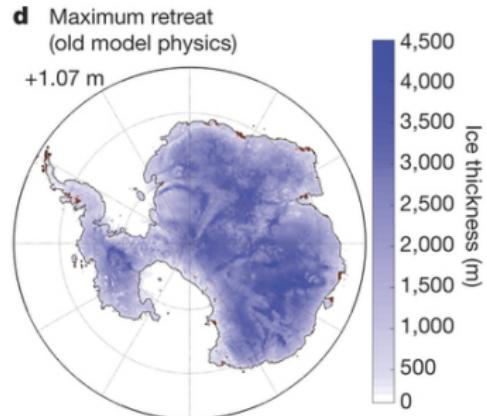
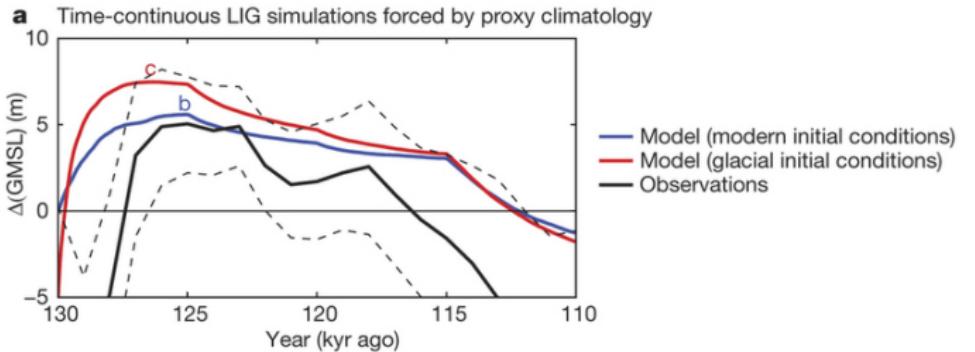
Claussen et al. (2002), *Climate Dynamics*

The development of coupled climate system models

Growth of Climate Modeling



Ice sheet modelling: The case of the ice that won't melt



DeConto and Pollard (2016), *Nature*

Ocean forcing: Changes in ocean circulation

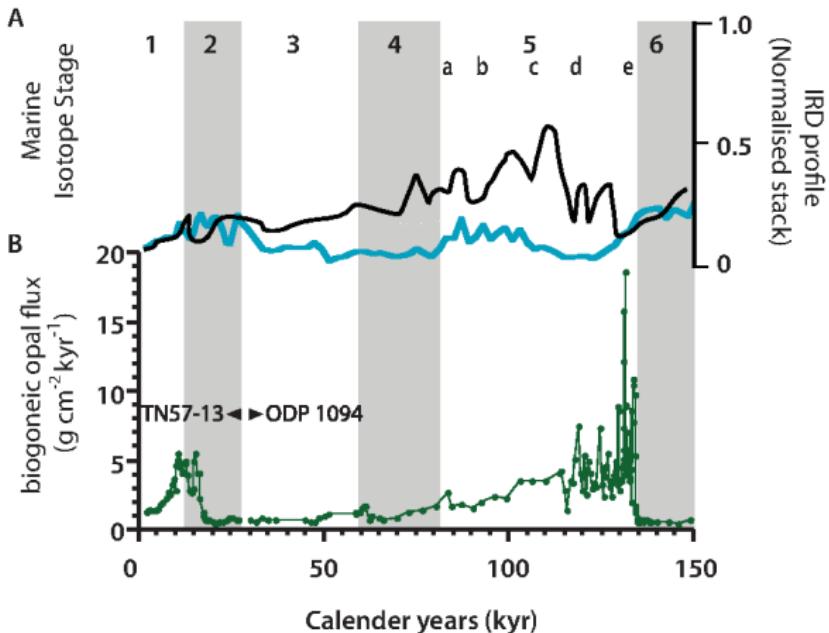
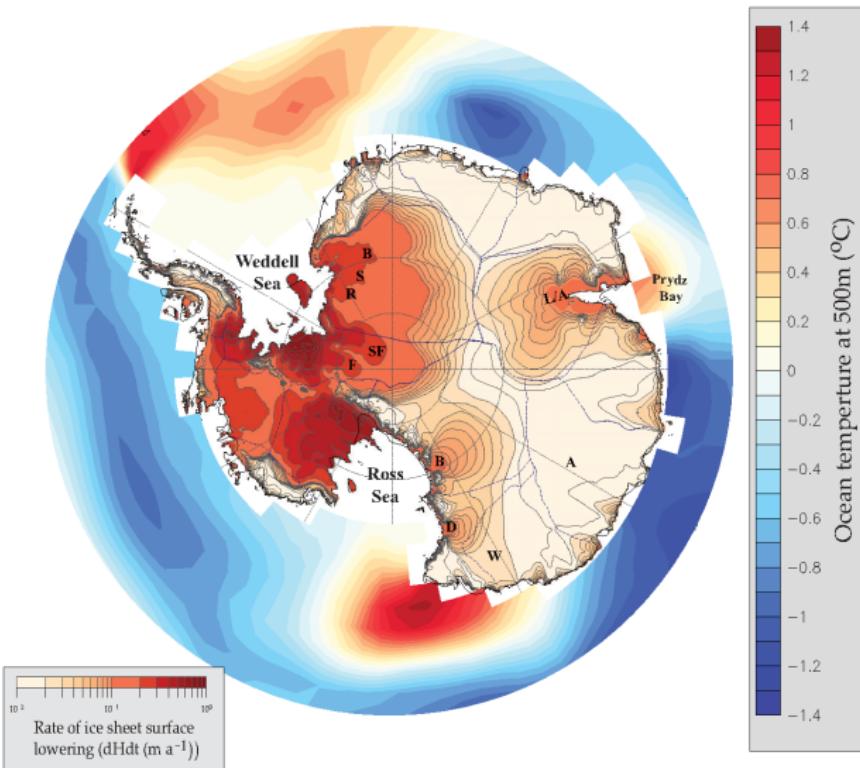


Figure 2. Proxy records of: (A) normalized stacked ice-rafter debris (IRD) profiles from the Weddell Sea (black line), and the Campbell Plateau (blue line). These composite records are based on datasets that were normalized to between zero and 1 (x-axis) by division of the maximum IRD concentration (Carter et al., 2002). (B) Biogenic particle flux, reconstructed by ^{230}Th normalization for biogenic opal flux, a proxy for Southern Ocean upwelling triggered by SHW southward migration (Jaccard et al., 2013), plotted against Marine Isotope Stages (MIS) over the past 150 ka. This figure is available in colour online at wileyonlinelibrary.com.

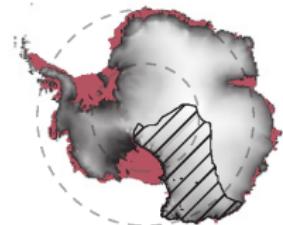
Fogwill et al. (2014), *Journal of Quaternary Science*

Ocean forcing: Changes in ocean circulation



Fogwill et al. (2014), *Journal of Quaternary Science*

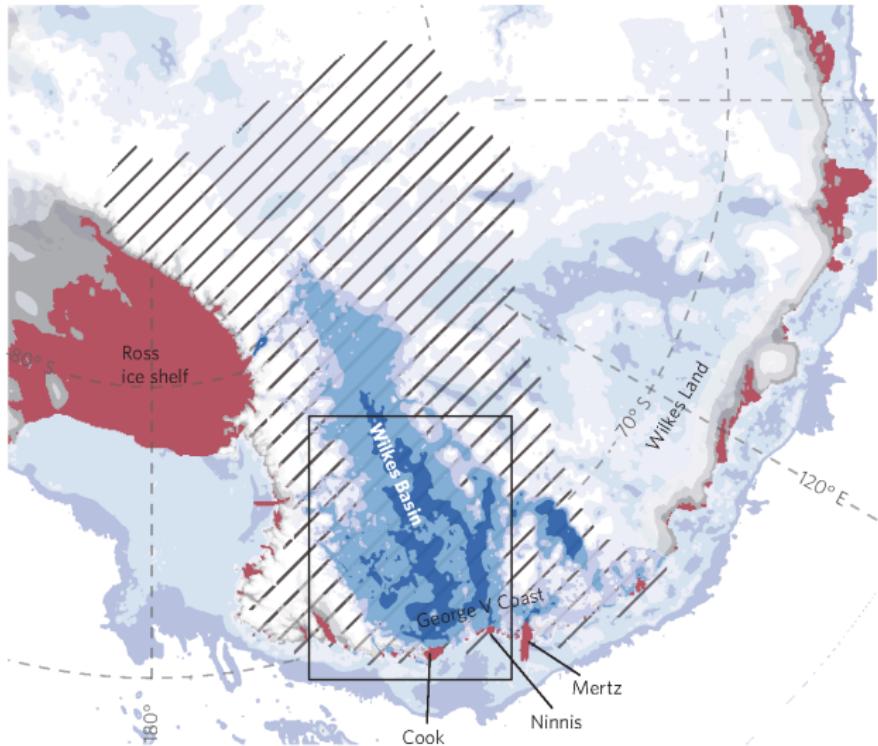
Ocean forcing: Marine ice sheet instability



Wilkes Basin
subglacial topography

- <1,000 m
- <500 m
- <0 m

Model domain



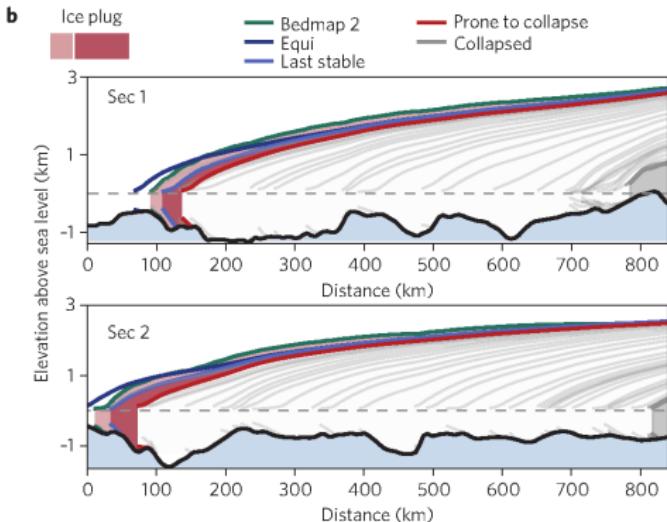
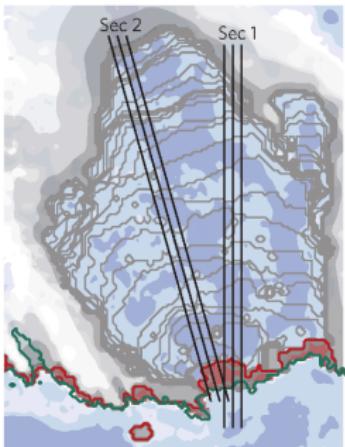
Mengel and Levermann (2014), *Nature Climate Change*

Ocean forcing: Marine ice sheet instability

a Subglacial topography
-<1,000 m
-<500 m
-<0 m

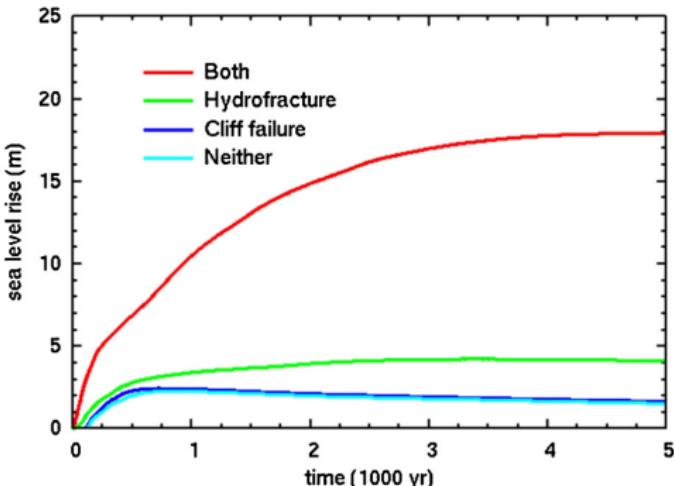
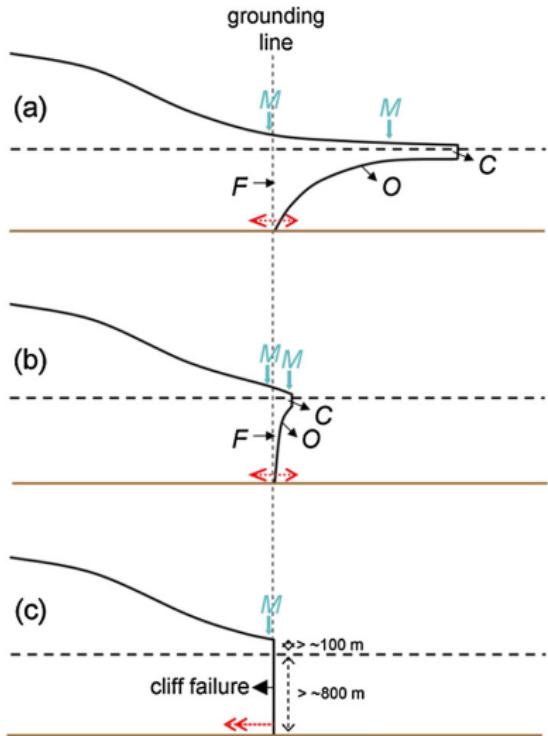
Ice thickness after retreat
<500 m
<1,000 m
<1,500 m

Grounding line
Bedmap 2
Prone to collapse
Transient retreat



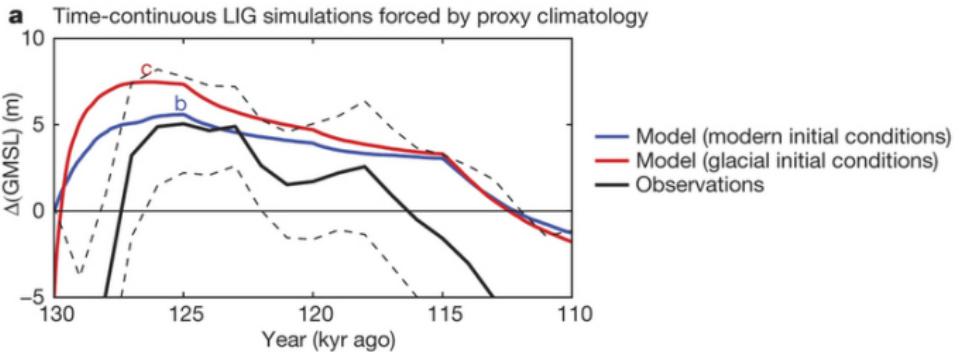
Mengel and Levermann (2014), *Nature Climate Change*

Ocean forcing: Hydrofracturing and ice cliff failure

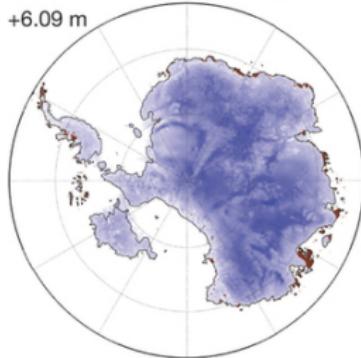


Pollard et al. (2015), *Earth and Planetary Science Letters*

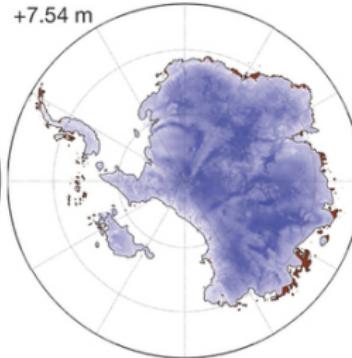
Ocean forcing: Marine ice cliff instability



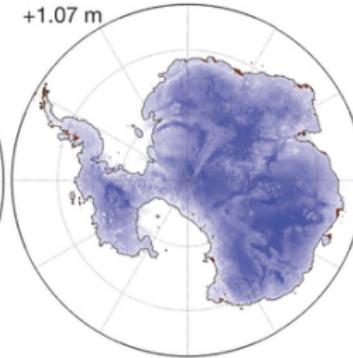
b Maximum retreat
(modern initial conditions)



c Maximum retreat
(glacial initial conditions)

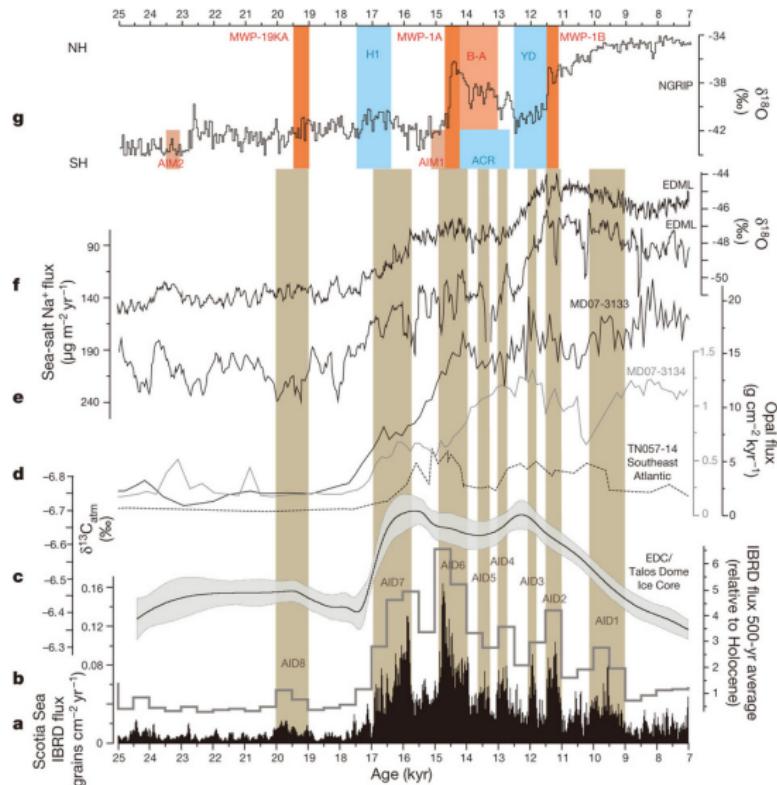


d Maximum retreat
(old model physics)



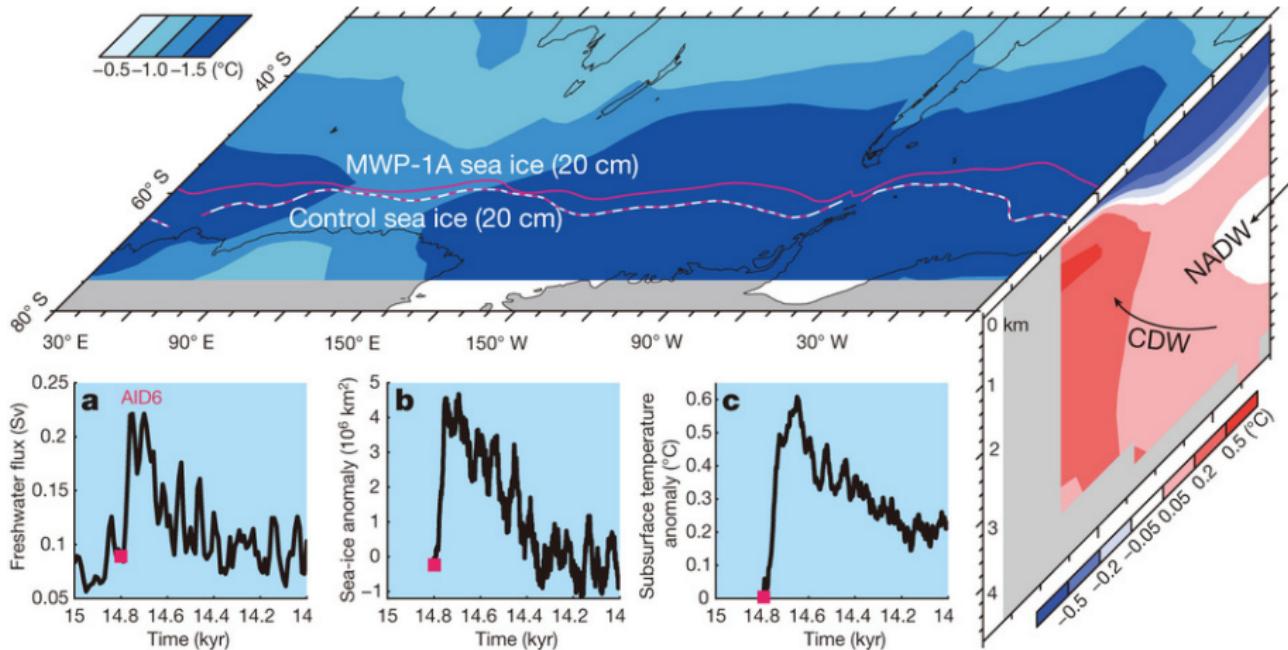
DeConto and Pollard (2016), *Nature*

Ocean forcing: Millennial-scale variability



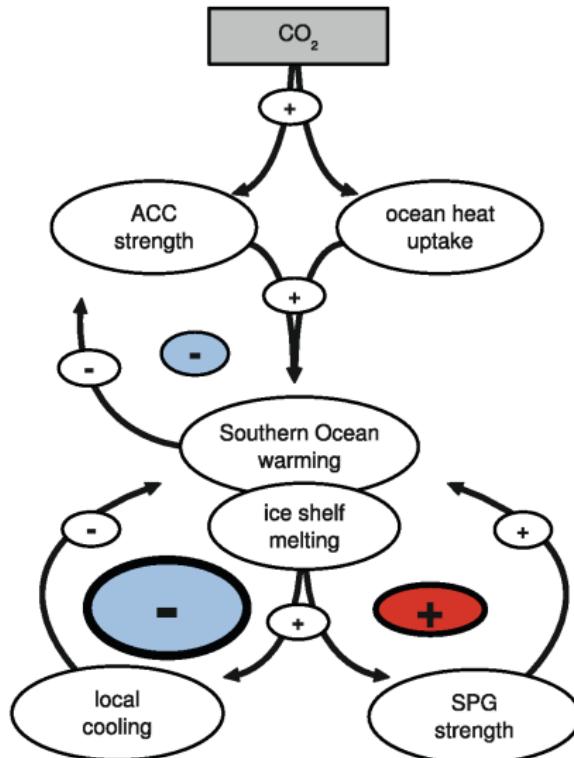
Weber et al. (2014), *Nature*

Ocean forcing: Millennial-scale variability



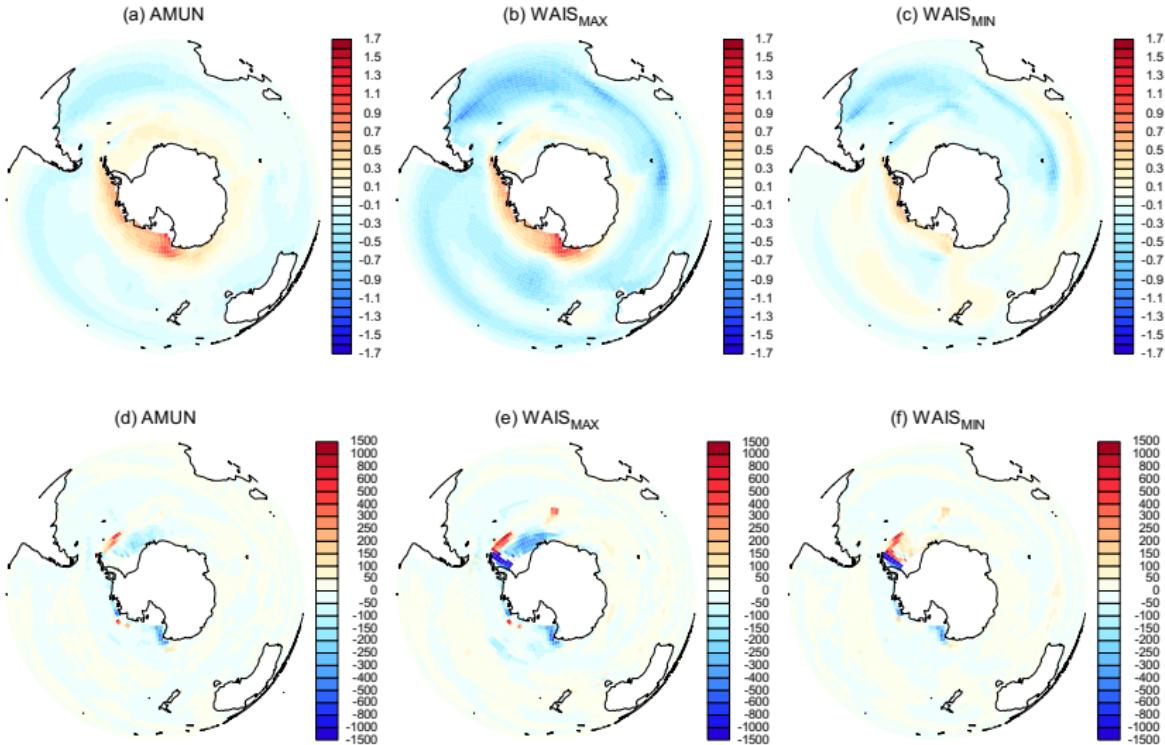
Weber et al. (2014), *Nature*

Ice sheet forcing: ice sheet–ocean feedback loops



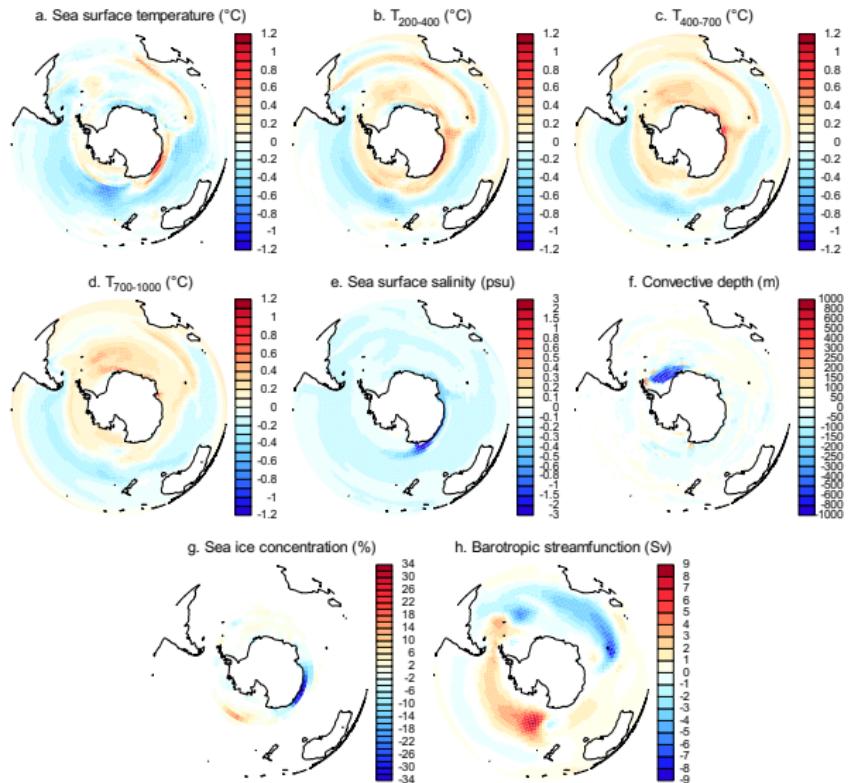
Hattermann and Levermann (2010), *Climate Dynamics*

Ice sheet forcing: freshwater fluxes into the ocean



Fogwill et al. (2015), *Earth's Future*

Ice sheet forcing: freshwater fluxes into the ocean



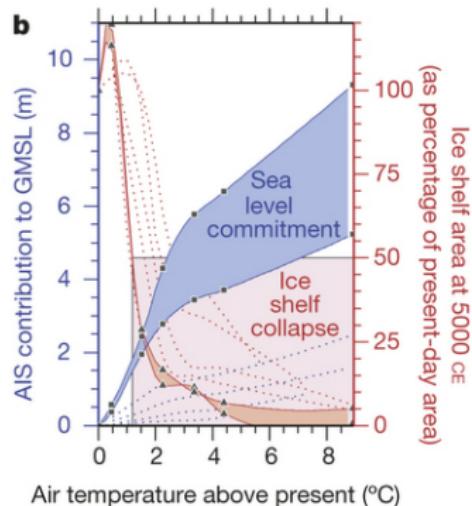
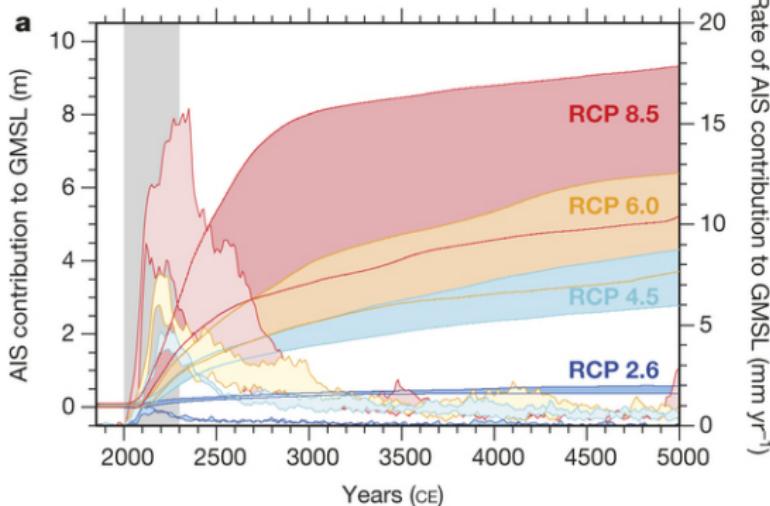
Phipps et al. (2016), *The Cryosphere*

Consequences for global sea level



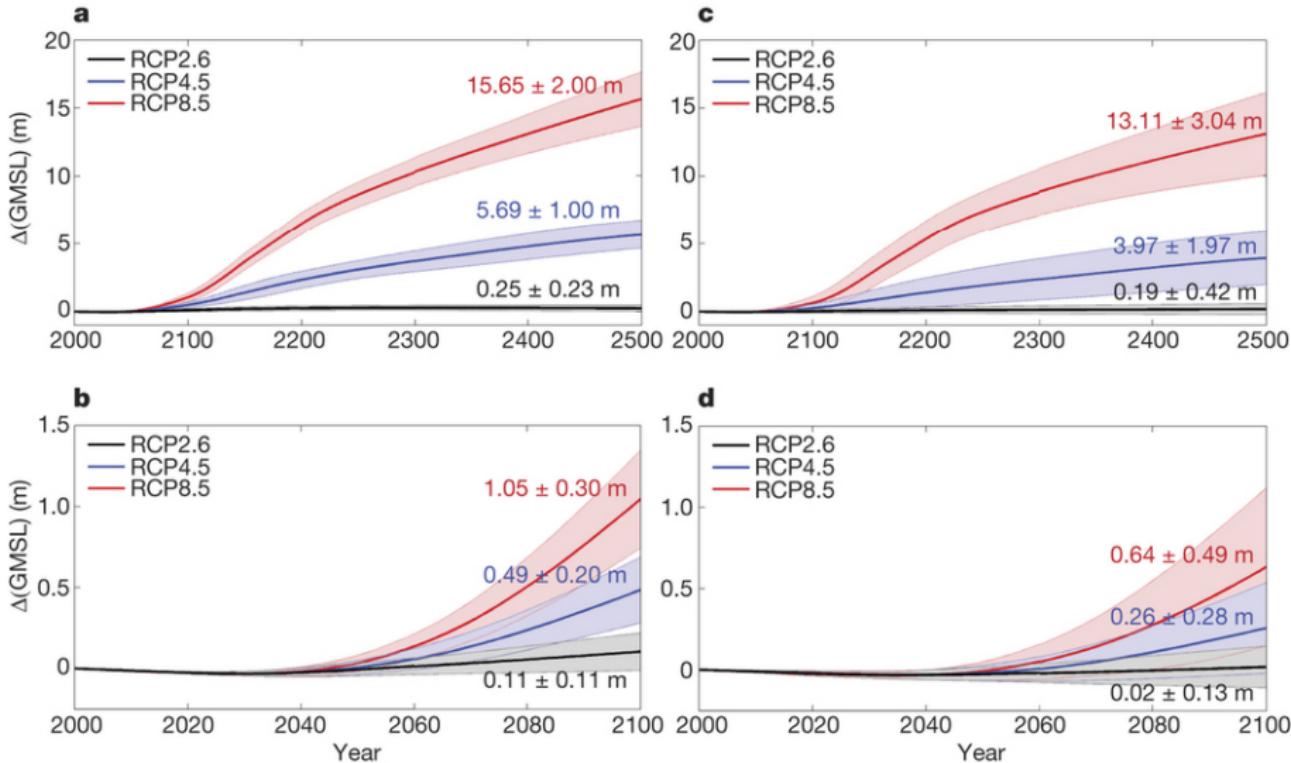
Golledge et al. (2015), *Nature*

Consequences for global sea level



Golledge et al. (2015), *Nature*

Consequences for global sea level



DeConto and Pollard (2016), *Nature*

Conclusions

- Ice sheet models can be used to explore the mechanisms that have driven past changes in the Antarctic ice sheet.
- Ice sheet models are the *only* physically-based tool that we have to predict future changes in the Antarctic ice sheet.
- Comparison with palaeoclimate data can be used to refine the models and identify potential missing physics.

Future research priorities:

- Continue to use palaeoclimate data to identify missing processes and incorporate them into ice sheet models.
- Couple, couple, couple...